# Calculations for ambient dose equivalent rates in nine forests in eastern Japan from <sup>134</sup>Cs and <sup>137</sup>Cs radioactivity measurements

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18	Highlights
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19	• Monte Carlo simulations for $\dot{H}^*(10)$ in forests in Japan between 2011 and 2017 from
20	<sup>134</sup> Cs and <sup>137</sup> Cs radioactivity measurements
21	• Mean residual error between simulated $\dot{H}^*(10)$ at 1 m and handheld survey meter
22	measurements was 16%.
23	• Calculated the contributions from <sup>134</sup> Cs and <sup>137</sup> Cs located in different parts of the
24	forests to $\dot{H}^*(10)$ .
25	• Largest contributions were from <sup>134</sup> Cs and <sup>137</sup> Cs in the canopy, organic layer and top
26	5 cm of soil.
27	• Uncertainties were evaluated in simulation results based on measurement uncertainties

in model inputs.

#### 29 Abstract

Understanding the relationship between the distribution of radioactive <sup>134</sup>Cs and <sup>137</sup>Cs in 30 forests and ambient dose equivalent rates  $(\dot{H}^*(10))$  in the air is important for researching forests 31 in eastern Japan affected by the Fukushima Dai-ichi Nuclear Power Plant (FDNPP) accident. 32 This study used a large number of measurements from forest samples, including <sup>134</sup>Cs and <sup>137</sup>Cs 33 34 radioactivity concentrations, densities and moisture contents, to perform Monte Carlo radiation transport simulations for  $\dot{H}^*(10)$  between 2011 and 2017. Calculated  $\dot{H}^*(10)$  at 0.1 and 1 m 35 above the ground had mean residual errors of 19% and 16%, respectively, from measurements 36 taken with handheld NaI(Tl) scintillator survey meters. Setting aside the contributions from 37 natural background radiation, <sup>134</sup>Cs and <sup>137</sup>Cs in the organic layer and the top 5 cm of forest soil 38 generally made the largest contributions to calculated  $\dot{H}^{*}(10)$ . The contributions from <sup>134</sup>Cs 39 and <sup>137</sup>Cs in the forest canopy were calculated to be largest in the first two years following the 40 accident. Uncertainties were evaluated in the simulation results due to the measurement 41 uncertainties in the model inputs by assuming Gaussian measurement errors. The mean 42 uncertainty (relative standard deviation) of the simulated  $\dot{H}^{*}(10)$  at 1 m height was 11%. The 43 main contributors to the total uncertainty in the simulation results were the accuracies to which 44 the <sup>134</sup>Cs and <sup>137</sup>Cs radioactivities of the organic layer and top 5 cm of soil, and the vertical 45 distribution of <sup>134</sup>Cs and <sup>137</sup>Cs within the 5 cm soil layers, were known. Radioactive cesium 46 47 located in the top 5 cm of soil was the main contributor to  $\dot{H}^*(10)$  at 1 m by 2016 or 2017 in the calculation results for all sites. Studies on the <sup>137</sup>Cs distribution within forest soil will 48 therefore help explain radiation levels henceforth in forests affected by the FDNPP accident. 49 The merits of this study are that it modelled multiple forests for a long time period, with the 50 important model inputs being informed by field measurements, and it quantified how the 51 measurement uncertainties in these inputs affected the calculation results. 52

- 54 Keywords: forests; ambient dose equivalent rate; simulation; Fukushima Dai-ichi Nuclear
- 55 Power Plant accident; Monte Carlo

## 56 **1. Introduction**

After the Fukushima Dai-ichi Nuclear Power Plant (FDNPP) accident in 2011, around 57 70% of the atmospheric <sup>134</sup>Cs and <sup>137</sup>Cs fallout in eastern Japan landed on forests (Hashimoto 58 et al., 2012; Kato et al., 2019). The management of these forests is an important issue in affected 59 regions as around 40% by area are plantations (Forestry Agency of Japan, 2012) and forestry 60 products make an important contribution to local economies (Miura, 2016). It has been shown 61 that <sup>134</sup>Cs and <sup>137</sup>Cs transfers slowly from forests to rivers (Evrard et al., 2015) and freshwater 62 fish (Wada et al., 2019; Kurikami et al., 2019). Thus economic and environmental 63 considerations need to be taken into account when evaluating future management choices for 64 the forests. 65

Studies often utilize mathematical relationships that connect ambient radiation levels 66 measured in the air with quantities of radioactive cesium for monitoring radioactivity levels in 67 68 the environment (UNSCEAR, 2000; Sanada et al., 2014; Mikami et al., 2015a; Karadeniz et al., 2015; Sakuma et al., 2018a). The numerical parameters in these relationships are derived from 69 70 radiation transport calculations (Beck et al., 1972; ICRU, 1994; Saito and Jacob, 1995; Saito and Petoussi-Henss, 2014). Calculations applicable for forests were first performed for sites 71 affected by atmospheric nuclear weapons test and Chernobyl accident fallout (Miller et al., 72 73 1990; Clouvas et al., 1999; Golikov et al., 1999; Gering et al., 2002; Clouvas et al., 2005). The relationship between <sup>134</sup>Cs and <sup>137</sup>Cs radioactivities in forests affected by the FDNPP accident 74 and radiation levels in the air has been modelled with various methodologies. These include 75 using simple analytical expressions for attenuation (Imamura et al., 2017a), or by interpolating 76 published <sup>134</sup>Cs and <sup>137</sup>Cs radioactivity to ambient dose equivalent rate conversion coefficients 77 to model forest soil layers and the organic layer (Teramage and Sun; 2020). More 78 comprehensive assessments have employed Monte Carlo radiation transport simulations 79 (Gonze et al., 2016; Hemmi, 2017; Cresswell et al., 2018: Sakuma et al., 2018b). 80

Radiation transport models of forests are necessarily more complex than models for open 81 82 sites, such as farmland and recreational spaces, as they need to account for the radioactivity present in trees and the forest organic layer, and the shielding of gamma rays within these parts 83 of forests. The radioactivity distribution in forests also changes dynamically over time due the 84 forest ecosystem (Hashimoto et al., 2020). Required input data for forest radiation transport 85 models include physical dimensions of trees, <sup>134</sup>Cs and <sup>137</sup>Cs radioactivity distributions, and 86 densities and moisture contents of different materials in forests. The large number of inputs 87 required mean that it is unusual for a single published forest measurement survey to contain all 88 necessary data for creating a realistic radiation transport model of a forest. 89

This study modelled ambient dose equivalent rates  $(\dot{H}^*(10))$  to high accuracy in nine 90 forests in eastern Japan for a long time period (2011–17) by collating measurements from 91 multiple field surveys to set up the models. The various measurements were used to simulate 92  $\dot{H}^*(10)$  at 0.1, 1 and 20 m above the ground using the Monte Carlo radiation transport method. 93 In addition to the main calculations for  $\dot{H}^*(10)$ , this paper presents the following: (i) 94 verification of the simulated  $\dot{H}^{*}(10)$  by comparing against measurements taken with handheld 95 NaI(Tl) scintillator survey meters in the forests; (ii) calculations of the separate contributions 96 to  $\dot{H}^*(10)$  from <sup>134</sup>Cs and <sup>137</sup>Cs located within the canopy, tree trunks, the forest organic layer 97 and different depth soil layers to determine the main sources of radioactivity contributing to 98  $\dot{H}^{*}(10)$ ; and (iii) a Monte Carlo uncertainty analysis of the effects of measurement uncertainties 99 in the model inputs on the simulation results. The merits of this study are that it modelled 100 multiple forests for a long period following the FDNPP accident in 2011, and that all the major 101 102 input parameters in the models were set based on survey data from forests in eastern Japan.

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# 104 **2. Methods**

# 105 2.1 Surveys by Forestry and Forest Products Research Institute

Physical measurements and <sup>134</sup>Cs and <sup>137</sup>Cs distributions for nine forests were surveyed by 106 the Forestry and Forest Products Research Institute (FFPRI) and the Forestry Agency of Japan 107 between 2011 and 2017 (Imamura et al., 2017b; Forestry Agency of Japan, 2017; 2018). Each 108 109 forest was referenced by a code denoting the forest location and the dominant species of trees. The forests were located in the Kawauchi (location code KU), Otama (OT), and Tadami (TD) 110 districts in Fukushima Prefecture, and the Ishioka (*Tsukuba* – TB) district in Ibaraki Prefecture. 111 The main tree species were evergreen coniferous trees of Japanese cedar (sugi, Cryptomeria 112 *japonica* – species code S), Japanese cypress (hinoki, *Chamaecyparis obtusa* – H), Japanese 113 red pine (akamatsu, Pinus densiflora – P) and deciduous broadleaf trees of konara oak (Quercus 114 serrata – Q). 115

Site OT-S is a single species plantation of Japanese cedar (Table 1). Sites KU1-S and TD-116 S contain mainly Japanese cedars, but also include some deciduous broadleaf species trees with 117 lower prevalence than Japanese cedar such as konara oak and deciduous Japanese chestnut 118 (Castanea crenata). Site KU2-S contains Japanese cedars, with deciduous conifers of Japanese 119 larch (karamatsu, Larix kaempferi) and other deciduous broadleaf species trees at lower 120 prevalence. Sites KU-H and TB-H are single species plantations of Japanese cypress. Site KU1-121 Q is a secondary forest of konara oaks and other broadleaf species trees. Sites OT-P and OT-Q 122 123 are secondary forests containing Japanese red pines and konara oaks plus other deciduous broadleaf species trees. 124

Research plots were established at each site with areas between 0.06 and 0.24 ha. Surveys were undertaken annually, usually in August or September, for the survey periods in Table 1. All measured <sup>134</sup>Cs and <sup>137</sup>Cs radioactivities in the samples were decay corrected to 1<sup>st</sup> September of the applicable sampling year.

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#### 130 **2.1.1 Tree measurements**

131 Stand counts and diameters at breast height (DBH) were taken for all trees (n = 4 to 268) 132 within each species group at the research plots (Kajimoto et al., 2014). Tree heights (m) were 133 estimated using DBH allometric equations applicable for each species group. The parameters 134 in the DBH allometric equations were derived using a sample of tree heights of the main species 135 trees measured using a laser rangefinder (Kajimoto et al., 2014).

Total oven-dried biomass quantities (kg/m<sup>2</sup>) of leaves/needles, branches, bark, sapwood 136 137 and heartwood for each species group were calculated using stand densities and DBH allometric equations. The parameters in DBH allometric equation were derived from samples taken from 138 main species trees that were felled each year from 2011 to 2016 (Kajimoto et al., 2014; Komatsu 139 140 et al., 2016). Trees (n = 3) with differing trunk diameters (small, medium and large) were felled each year during this period (Kuroda et al., 2013; Kajimoto et al., 2014). Wood disks were cut 141 from the felled trunks and split into bark, sapwood and heartwood sections. For 2017, samples 142 from trunks were taken by the tree ring core method at sites KU1-H, KU1-Q, KU1-S and KU2-143 S, while three trees were felled at site OT-Q as per the sampling methodology in earlier years. 144 145 Branch and needle/leaf samples were taken for 2011 to 2017 (n = 3 to 4).

The samples were oven-dried, weighed and the <sup>134</sup>Cs and <sup>137</sup>Cs contents (Bq/kg) measured using gamma spectroscopy (n = 3 to 4) (Imamura et al., 2017b; Kenzo et al., 2020). Lowest live branch heights were measured on the sampled trunks (n = 3). Moisture contents (defined as the ratio of the water to oven-dried material mass) of the bark, sapwood and heartwood were calculated from the wood disk samples (n = 4 to 8) (Kuroda et al., 2013; Ohashi et al., 2017).

For forests with more than one group of tree species (sites KU1-S, KU2-S, OT-Q, OT-P and TD-S), data (including DBH equations) for the lower prevalence group(s) were inferred using data from another site where the corresponding species group was the main species group (Table S1). For example, data for Japanese red pines at site OT-Q were estimated using the Japanese red pine data from site OT-P. For the cases where it was necessary to infer data from other sites, <sup>134</sup>Cs and <sup>137</sup>Cs radioactivities were scaled to account for any intrinsic differences
in fallout radioactivity levels between the sites (Komatsu et al., 2016).

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# 159 2.1.2 Forest organic layer and soil measurements

Samples of the organic layer were gathered from  $25 \times 25$  cm quadrats on the forest floor 160 (n = 6 to 12). Soil was sampled in 5 cm layers down to 20 cm depth using a metal core sampler 161 (*n* = 3 to 12) (Fujii et al., 2014; Ikeda et al., 2014; Komatsu et al., 2016; Manaka et al., 2019). 162 The samples were oven-dried, weighed and the <sup>134</sup>Cs and <sup>137</sup>Cs radioactivities measured using 163 high purity gamma spectroscopy detectors. The results were used to calculate mean oven-dried 164 densities (kg/m<sup>3</sup>) and <sup>134</sup>Cs and <sup>137</sup>Cs inventories (Bq/m<sup>2</sup>) of the organic layer and soil layers. 165 Following Imamura et al. (2017a), the organic layer was assumed to be 4 cm thick for the 166 calculations. 167

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# 169 2.1.3 Measurement of ambient dose equivalent rates

Ambient dose equivalent rates at 0.1 and 1 m heights above the ground surface were measured using handheld NaI(Tl) scintillator survey meters (TCS-172B, Hitachi-Aloka Medical, Ltd., Tokyo, Japan). Measurements were taken at the vertices of  $10 \times 10$  m grid squares within each plot (Ikeda et al., 2014; Imamura et at., 2015; Komatsu et al., 2016). The mean values of the measurements (n = 6 to 36) at each height were used for comparison with the simulation results.

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#### 177 **2.2 Data from surveys of other forest sites**

178 Not all measurements necessary for creating a radiation transport model to link the <sup>134</sup>Cs 179 and <sup>137</sup>Cs radioactivity distribution in the forests with  $\dot{H}^*(10)$  were available from the surveys 180 by FFPRI. The moisture contents of the branches and needles/leaves of trees, the organic layer

and the soil layers were not measured during the FFPRI surveys, but these data are needed to 181 model the overall densities and elemental compositions of these materials, which affect the 182 attenuation of <sup>134</sup>Cs and <sup>137</sup>Cs gamma rays. The vertical distribution of <sup>134</sup>Cs and <sup>137</sup>Cs in soil 183 was also only measured to the resolution of 5 cm thick soil layers in these surveys. In reality it 184 is typical for the concentration of <sup>134</sup>Cs and <sup>137</sup>Cs in forest soils in eastern Japan to decrease 185 exponentially with depth (Teramage et al., 2016; Yoschenko et al., 2017; Muto et al., 2019). 186 Previous studies have reported a relaxation depth of the exponential function,  $\alpha$ , on the order of 187 0.5 to 5 cm (Mishra et al., 2016; Takahashi et al., 2018). Thus it is necessary to model the 188 vertical distribution of <sup>134</sup>Cs and <sup>137</sup>Cs in soil with a finer resolution than 5 cm thick soil layers 189 for forest  $\dot{H}^{*}(10)$  calculations. As outlined below, moisture contents for the canopy, organic 190 layer and soil, and data on the vertical distribution of <sup>134</sup>Cs and <sup>137</sup>Cs within soil, were obtained 191 from different forest surveys for set up the simulation models. 192

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#### **194 2.2.1 Moisture contents of branches and needles/leaves**

Data for the moisture contents of branches and needles/leaves came from forest sites in 195 the Kawamata and Kawauchi districts of Fukushima Prefecture monitored by Japan Atomic 196 Energy Agency (JAEA) (Niizato et al., 2016). Needle and branch samples of Japanese cedars 197 were taken from the upper, middle and lower parts of the crowns of five trees in the Kawauchi 198 forest on 27–29 October 2015 and on 25–27 September 2017. The mean moisture contents of 199 200 these samples were used to model the moisture contents of the needles and branches of the evergreen coniferous trees at the sites monitored by FFPRI. The mean moisture contents of 201 leaves and branches sampled from the crowns of five konara oak trees in a forest in Kawamata 202 203 on 26-29 November 2018 were used to model the moisture contents of deciduous broadleaf tree leaves and branches at the sites monitored by FFPRI. 204

#### 206 2.2.2 Moisture contents for the organic layer and soil

Moisture contents for the organic layer and soil layers were also based on data from forests in Kawamata and Kawauchi surveyed by JAEA between 2015 and 2017. Samples of the organic layer and soil were taken from Japanese cedar forests in Kawauchi in the periods 1–3 December 2015, 18–20 October 2016, 21–23 August 2017 and 16–20 October 2017. Means (n = 10) of the moisture content of the organic layer and 5 cm thick soil layers down to 20 cm depth were used to model the single stand evergreen coniferous species forests monitored by FFPRI (KU1-H, OT-S, TB-H).

Data from forests containing deciduous broadleaf species and Japanese red pine trees in Kawamata and Kawauchi surveyed by JAEA was used for modelling the moisture content of the organic layer and soil in the forests monitored by FFPRI with multiple tree species. These surveys took place on 3–8 December 2015, 17–21 October 2016 and 17–31 October 2017. Means (n = 7) of the moisture contents were used in the simulations.

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# 220 **2.2.3** Vertical distribution of <sup>134</sup>Cs and <sup>137</sup>Cs within 5 cm soil layers

The vertical distribution of <sup>137</sup>Cs within soil at three forest sites at Yamakiya, Kawamata was monitored by the University of Tsukuba between 2011 and 2017 (Takahashi et al., 2018). The three sites are a mixed forest (MF) of Japanese red pine and konara oak, and a young cedar (YC) and mature cedar (MC) plantation of Japanese cedar. A scraper plate (450 cm<sup>2</sup>) area was used to collect soil layers from each site over nine surveys between June 2011 and June 2017. The data all show an approximately exponential decrease in the <sup>137</sup>Cs concentration with vertical depth in soil.

Following Liu et al. (2019), the cumulative form of the exponential distribution

$$I(\zeta) = I_{inv}(1 - \exp(-\zeta/\beta)) \tag{1}$$

was used to fit the cumulative inventory of <sup>137</sup>Cs,  $I(\zeta)$  (Bq/cm<sup>2</sup>), as a function of the cumulative mass depth of wet soil,  $\zeta$  (g/cm<sup>2</sup>). Relaxation mass depth parameters,  $\beta$  (g/cm<sup>2</sup>), and total <sup>137</sup>Cs inventories,  $I_{inv}$  (Bq/cm<sup>2</sup>), were extracted from the fits. The fitting was performed using the trust region reflective algorithm for non-linear least squares optimization in the SciPy Python package (Virtanen et al., 2020). Fitted parameters are listed in Table S2 and fits shown in Fig. S1.

236 Only one set of scraper plate samples was taken at each site per survey and there was a large degree of scatter in the fitted  $\beta$  values between different survey dates (Fig. 1). It is not 237 clear a priori that  $\beta$  should increase monotonically over time in forest soils as is the case for 238 open sites (Liu et al., 2019). This is because at open sites the input of <sup>134</sup>Cs and <sup>137</sup>Cs to the soil 239 surface occurred acutely in March 2011 with the radioactive plumes from the FDNPP accident. 240 For forests, the input of <sup>134</sup>Cs and <sup>137</sup>Cs into forest soil occurs over a prolonged period, starting 241 with the initial deposition fallout, followed by the transfer of <sup>134</sup>Cs and <sup>137</sup>Cs from the forest 242 canopy to the organic layer and from the organic layer to soil over many years affected by 243 244 multiple processes. Given the scatter of the data in Fig. 1 explained by the low number of samples, it was not considered justifiable to fit a linear function for the dependency of  $\beta$  on time 245 to the data. 246

Instead, the mean values of  $\beta$  from the surveys of the YC and MC sites (n = 18), and from the surveys at the MF site (n = 9), were used to model the distribution of <sup>134</sup>Cs and <sup>137</sup>Cs within the 5 cm layers at the single stand and multiple stand forests monitored by FFPRI, respectively. It is recognized that this method means the simulations cannot model how temporal changes to the <sup>134</sup>Cs and <sup>137</sup>Cs mass depth profiles inside the 5 cm forest soil layers affect  $\dot{H}^*(10)$ . The potential effects of this limitation were therefore assessed with the uncertainty analysis. Note however that the migration of <sup>134</sup>Cs and <sup>137</sup>Cs between the organic layer and different 5 cm soil layers is captured by the simulations, as the FFPRI surveys measured the radioactivities of thesedifferent layers independently.

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# 257 2.2.4 Natural background radiation

Results from Andoh et al. (2017) were used to estimate the natural background component 258 in the  $\dot{H}^*(10)$  measurements taken with NaI(Tl) scintillator survey meters (TCS-172B) during 259 260 the FFRPI surveys. Andoh et al. (2017) analyzed gamma spectra measured with KURAMA-II instruments (CsI(Tl) scintillators) (Tanigaki et al., 2015) during car-borne surveys in eastern 261 Japan to separate the natural background component of  $\dot{H}^{*}(10)$  from the <sup>134</sup>Cs and <sup>137</sup>Cs 262 components. The mean natural background component of  $\dot{H}^{*}(10)$  calculated for Fukushima 263 Prefecture of 0.055 µSv/h (±0.010 µSv/h) was used for modelling the natural background 264 component for the sites in Fukushima Prefecture. The value calculated for Ibaraki Prefecture 265  $(0.060\pm0.006 \,\mu\text{Sv/h})$  was used for site TB-H. Note that the natural background component of 266  $\dot{H}^{*}(10)$  in the measurements taken with the TCS-172B and KURAMA-II survey meters 267 includes contributions from both terrestrial gamma rays and cosmic radiation. The terrestrial 268 269 component is the dominant component however, as these devices have low sensitivity to cosmic 270 radiation (Nagaoka et al., 1996; Yajima et al., 2014; Andoh et al., 2017).

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#### 272 **2.3 Radiation transport simulations**

The simulations for  $\dot{H}^*(10)$  were undertaken with the Particle and Heavy Ion Transport code System (PHITS) version 3.20, which is a Monte Carlo code (Sato et al., 2018). A model of multiple layers of uniform matter was used to represent forests (Fig. 2). The thickness, density, material composition, and <sup>134</sup>Cs and <sup>137</sup>Cs radioactivities were set independently for each layer based on the measurements from the various forest samples. The models were created in boxes with reflective boundary conditions for gamma rays at the vertical boundary faces. This method ensured an effectively infinite area forest was simulated so that the  $\dot{H}^*(10)$ estimates would saturate (Malins et al., 2015a; Malins et al., 2015b). It can be used because the model geometry and the <sup>134</sup>Cs and <sup>137</sup>Cs source distribution has 4-fold rotational symmetry.

282 The elemental compositions of the materials for each layer were set by combining base components of air, water, organic matter and mineral soil (Table 2). The density of each base 283 component was equal to its respective volumetric density within the layer. Soil was modelled 284 285 as a mixture of black forest soil (andosol), organic matter and water. The carbon content of dry forest soil (i.e. the black forest soil and organic matter mixture) was set to 10.7% by weight 286 287 (Kawada, 1982; Kawada, 1989). The organic layer was assumed to contain zero black forest soil. The canopy and trunk layers were modelled as a mixture of dry organic matter, water and 288 289 air.

The means of the physical data used for each layer in the forest models are summarized in Table 3. The masses per unit area include both the dry material and the moisture masses. For the trunk and canopy layers, the data include the air mass per unit area (calculated using a volumetric density of  $1.2 \text{ kg/m}^3$ ). Note that tree matter only displaces a negligible volume of the atmosphere. Maximum and minimum values of the data are given for the nine sites.

The Electron-Gamma Shower version 5 (EGS5) algorithm in PHITS was used for photon transport (Hirayama et al., 2005; Sato et al., 2018). The data for energies and emission probabilities for gamma rays from <sup>134</sup>Cs and <sup>137</sup>Cs decay came from the Evaluated Nuclear Structure Data Files (ENSDF). This data was accessed using the Nudat2 (2020) website.

Values of  $\dot{H}^*(10)$  were evaluated using track length tallies to calculate photon fluxes, then these were converted to  $\dot{H}^*(10)$  using conversion coefficients from ICRP (1996). Tally regions were thin layers set at 0.1, 1 and 20 m above the ground (dashed lines, Fig. 2). Monte Carlo statistical uncertainties for calculated  $\dot{H}^*(10)$  were lower than 0.1%. The minimum energy cut off for photon transport in EGS5 was set at 3 keV. For computational efficiency secondary particles from photon interactions (electrons, positrons and Bremsstrahlung) were not generated and transported in the simulations. The effect of this simplification was checked by repeating one simulation will full transport of all particles (photons, electrons, positrons and Bremsstrahlung) down to 1 keV. The difference in the simulated  $\dot{H}^*(10)$  between the two cases was negligible.

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# 310 2.4 Analysis of uncertainties

311 The uncertainties in the parameters derived from the field measurements that were used as simulation inputs were characterized as follows. Standard errors of measurements were used to 312 represent the uncertainties when the measurements were taken at the same site and in the same 313 year as the case being simulated. Standard deviations were used otherwise, i.e. for the scenarios 314 where specific measurements for that site and year were not available, so the data had to be 315 estimated using measurements taken in other years or from other sites. This is because the 316 plausible distributions of the unmeasured values are better represented by standard deviations 317 of the proxy measurements than standard errors. 318

The uncertainties in the simulation results for  $\dot{H}^*(10)$  that propagate from the uncertainties in the measured input variables were assessed as follows. It was assumed that the distributions of the measurement uncertainties were Gaussian. For each site and year, multiple simulations were performed with the input variables drawn randomly from the uncertainty distributions. For simplicity it was assumed that the uncertainties for each variable were uncorrelated. The exception to this was the <sup>134</sup>Cs and <sup>137</sup>Cs radioactivity measurements of each component, which were assumed to be perfectly correlated to keep the <sup>134</sup>Cs/<sup>137</sup>Cs ratio fixed.

Total uncertainties in the simulated  $\dot{H}^*(10)$  were calculated using the above method by varying all input variables randomly within their respective uncertainty distributions. To understand the contributions from different input uncertainties to the total uncertainty, further simulations were performed for sites OT-Q, OT-S and TB-H for 2011 and 2016 varying only a single class of input variables randomly within their uncertainty distributions. It was found that performing 100 simulations with randomly drawn input variables was sufficient to converge the uncertainty estimates for the simulated  $\dot{H}^*(10)$  (Fig. 3).

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#### **334 3. Results and discussion**

#### **335 3.1** Comparison of simulated and measured ambient dose equivalent rates

There was a clear correspondence between the simulated  $\dot{H}^*(10)$  and the measurements taken with handheld NaI(Tl) scintillator survey meters (Fig. 4). All simulation results were within a factor of two of the measurements. The mean absolute percentage deviations of the simulations from the measurements were respectively 24% and 22% for  $\dot{H}^*(10)$  at 0.1 and 1 m above the ground (Table 4).

The correlation between simulated and measured  $\dot{H}^*(10)$  in this study compares favorably 341 against a previous study simulating  $\dot{H}^*(10)$  at open sites, such as farmland, grass fields and 342 parks (Malins et al., 2016). In that study the mean absolute percentage deviation of the 343 simulations from measurements was 29%, and the coefficient of determination ( $R^2$ ) was 0.825. 344 The range of dose rates at sites in both studies was similar (0.05 to 5  $\mu$ Sv/h). As the modelling 345 346 methods and complexity were broadly comparable between the two studies, the better correlation obtained in this study is attributable to the lower uncertainties in inputs to the 347 simulations. The simulations in Malins et al. (2016) were based on a single scraper plate sample 348 set taken at each site. In comparison in this study the simulation inputs were typically averages 349 of n = 3 to 6 replicated measurements taken at each site per survey. 350

There was also better agreement between simulated and measured  $\dot{H}^*(10)$  in this study than a recent study which created 3D Monte Carlo models of sites in Fukushima Prefecture with individual buildings and trees tailored to specific sites (Kim et al., 2019). Again the likely
explanation for this is the lower measurement uncertainties in the inputs to the models in this
study due to more replications of measurements.

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# 357 **3.2 Breakdown of different source contributions to ambient dose equivalent rates**

Breakdowns of the contributions to the simulated  $\dot{H}^*(10)$  at 1 m above the ground from <sup>134</sup>Cs and <sup>137</sup>Cs located in different parts of the forests and from natural background radiation are shown in Fig. 5. The largest contributors to simulated  $\dot{H}^*(10)$  over the study periods were <sup>134</sup>Cs and <sup>137</sup>Cs located in the canopy, organic layer and top 5 cm of soil, as well as natural background radiation for sites with low overall <sup>134</sup>Cs and <sup>137</sup>Cs radioactivity levels like TB-H and TD-S.

Contributions from <sup>134</sup>Cs and <sup>137</sup>Cs in the canopy were greatest in 2011 and 2012, as the interception of atmospheric fallout by the canopy in March 2011 meant the radioactivity of the canopy was highest in the first two years (Fig. 6). The radioactivity in the canopy transferred to the forest floor due to throughfall, litterfall and stemflow in the years following the fallout in 2011 (Teramage et al., 2014; Endo et al., 2015; Kato et al., 2017; Imamura et al., 2017b). This meant the contributions from radioactivity in the canopy to  $\dot{H}^*(10)$  decreased substantially in later years of the study periods (Fig. 5).

There were large contributions to simulated  $\dot{H}^{*}(10)$  from <sup>134</sup>Cs and <sup>137</sup>Cs in the organic layer and top 5 cm of soil from 2011 onwards (Fig. 5). The transfer of <sup>134</sup>Cs and <sup>137</sup>Cs in the organic layer to the top 5 cm of soil occurs more gradually than from the canopy to the forest floor (Hashimoto et al., 2020), cf. Fig. 6. This meant that, setting aside the contribution from natural background radiation at the lower dose rate sites, there was a slow change to radioactive cesium within the top 5 cm of soil becoming the main contributor to  $\dot{H}^{*}(10)$  by 2016 and 2017 (Fig. 5).

The relative contribution of  ${}^{134}$ Cs and  ${}^{137}$ Cs in the organic layer to simulated  $\dot{H}^*(10)$ 378 becoming progressively smaller over time underlies why decontaminating forests by removing 379 the organic layer becomes less effective the later it is conducted (Koarashi et al., 2020). Unless 380 381 the organic layer is replaced by a fresh layer of uncontaminated material, e.g. wood chips (Cresswell et al., 2016), removing it in the later years of the 2011-17 study period will 382 undercover more highly radioactive soil beneath the organic layer. The loss of shielding of this 383 soil by the organic layer will then act counteractively to the benefit of removing the <sup>134</sup>Cs and 384 <sup>137</sup>Cs inventory associated with the organic layer from the forest. 385

Cesium-134 and <sup>137</sup>Cs located in tree trunks and in soil between 5 and 20 cm depths makes 386 a minor contribution to  $\dot{H}^*(10)$  for all sites and years simulated (Fig. 5). The former is 387 explained by the low quantity of <sup>134</sup>Cs and <sup>137</sup>Cs contained within trunks for all years (Fig. 6), 388 which is typically less than 3% of the total radioactivity of <sup>134</sup>Cs and <sup>137</sup>Cs in the forests 389 (Imamura et al., 2017b). For the latter, the quantity of <sup>134</sup>Cs and <sup>137</sup>Cs within soil between 5 and 390 20 cm depths only attains around 15% of the total forest <sup>134</sup>Cs and <sup>137</sup>Cs inventory by 2016 and 391 2017 (Fig. 6). The self-shielding of these gamma rays by the topsoil means  $^{134}$ Cs and  $^{137}$ Cs in 392 soil between 5 and 20 cm depths only contributes to 8–10% of the total  $\dot{H}^*(10)$  in 2016 and 393 2017. 394

The relative contribution form natural background radiation ranges from 3.8% (KU1-S, 2011) to 63% (TD-S, 2015) of the total  $\dot{H}^*(10)$  for the sites surveyed. In all cases the relative importance of the background contribution increases over time as radioactive decay reduces the quantities of <sup>134</sup>Cs and <sup>137</sup>Cs in the forests.

Fig. 7 shows the contributions to simulated  $\dot{H}^*(10)$  at 20 m above the ground. This height was chosen for analysis as it is above the upper canopy of all the forests and is a typical height for the operation of unmanned vehicles in aerial radiation surveys (Ochi et al., 2019). The relative importance of <sup>134</sup>Cs and <sup>137</sup>Cs within the canopy to  $\dot{H}^*(10)$  above it is larger than for 403  $\dot{H}^*(10)$  near ground level (cf. Figs. 5 and 7). The canopy contribution is on average 48% of the 404 total contribution from <sup>134</sup>Cs and <sup>137</sup>Cs to  $\dot{H}^*(10)$  at 20 m in 2011, decreasing to 8.9% by 2014. 405 The simulated  $\dot{H}^*(10)$  above the canopy decreased at a faster rate than at ground level 406 following fallout, consistent with modelling by Gonze et al. (2016). This effect may need to be 407 accounted for in the analysis of airborne survey results performed in different years following 408 radioactive fallout events (Kato et al., 2019).

409

# 410 **3.3 Effects of uncertainties in input measurements on simulations**

The mean relative uncertainty in simulated  $\dot{H}^*(10)$  at 1 m above the ground due to the measurement uncertainties in the model inputs was 11%. This was four times greater than the mean relative uncertainty, 2.8%, in the  $\dot{H}^*(10)$  measurements at 1 m. The breakdown of factors contributing to the overall simulation uncertainty is shown for three sites for 2011 and 2016 in Fig. 8 (right vertical axis).

The main contributors to the overall simulation uncertainties were the measurement uncertainties in the <sup>134</sup>Cs and <sup>137</sup>Cs radioactivity of the top 5 cm of soil, the organic layer and the canopy, the relaxation mass depth parameter  $\beta$  for the distribution of <sup>134</sup>Cs and <sup>137</sup>Cs in the 5 cm soil layers, and the natural background contribution to  $\dot{H}^*(10)$ . This reflects both the high relative uncertainties in the measurements underlying these inputs (shown by black dots and lines, left vertical axis in Fig. 8), and that the majority of the <sup>134</sup>Cs and <sup>137</sup>Cs in the forests was contained in these layers over the study period.

The uncertainty in the distribution of <sup>134</sup>Cs and <sup>137</sup>Cs with mass depth in the 5 cm soil layers has an increasingly important effect as <sup>134</sup>Cs and <sup>137</sup>Cs enters the topsoil in the years following the fallout. In this study the relaxation mass depth parameter ( $\beta$ ) was fixed for each site for all simulated years. By consequence of linking the vertical distribution of <sup>134</sup>Cs and <sup>137</sup>Cs within each 5 cm soil layer to  $\beta$ , the input uncertainties in the dry density and moisture 428 content of the soil layers had negligible effect on the simulated  $\dot{H}^*(10)$  results. This is because 429 the <sup>134</sup>Cs and <sup>137</sup>Cs mass depth distribution that is most important factor determining the gamma 430 ray fluence in the air (ICRU, 1994) and keeping  $\beta$  fixed means the <sup>134</sup>Cs and <sup>137</sup>Cs mass depth 431 distribution in soil remains unchanged if the soil dry density or moisture content are varied.

In reality the vertical distribution of <sup>134</sup>Cs and <sup>137</sup>Cs within forest soil varies on the 432 centimeter scale over time (Muto et al., 2019) and this potentially affects ambient dose 433 equivalent rates in forests in a similar manner to open sites (Mikami et al., 2015b). 434 Understanding the importance of this effect on  $\dot{H}^*(10)$  in forests remains an open question that 435 will require better data on the temporal dependency of the vertical distribution of <sup>134</sup>Cs and 436 <sup>137</sup>Cs in forests soil over time to resolve. However, the results of Fig. 8 suggest that variations 437 in the relaxation mass depth parameter  $\beta$  between approximately 0.5 to 2 g/cm<sup>2</sup> in forest soils 438 affects  $\dot{H}^{*}(10)$  on the order of 5%. On a timescale of decades, the radioactive cesium 439 440 distribution in forest soil may tend towards a more uniform distribution due to root decomposition (Sakashita et al., 2020). Monitoring studies of the radioactive cesium 441 442 distribution in forest soil will therefore be important over the long term.

The uncertainty in the natural background contribution to  $\dot{H}^*(10)$  at these sites (Fig. 8) could be reduced by performing in situ gamma spectroscopy measurements at the sites and calculating the natural background contribution from the measured spectrum (Clouvas et al., 1999; Cresswell et al., 2016). This would reduce the uncertainties that arose from using the data from car-borne surveys aggregated at district and prefectural levels (Andoh et al., 2017) to infer this component for these models.

449

# 450 **4. Conclusions**

451 Monte Carlo radiation transport simulations were performed for ambient dose equivalent 452 rates in nine forests in eastern Japan for 2011 to 2017. The simulation inputs were set using measurements from forests in eastern Japan. The measurement data used included dry densities,
 moisture contents and <sup>134</sup>Cs and <sup>137</sup>Cs inventories of different forest compartments.

There was a good correspondence between simulated  $\dot{H}^*(10)$  and NaI(Tl) scintillator survey meter measurements. The correlation between simulated and measured  $\dot{H}^*(10)$  was better in this study on forests than in previous studies modelling  $\dot{H}^*(10)$  at open sites and suburban locations (Malins et al., 2016; Kim et al., 2019). This was attributed to more replications of measurements underlying the input data in the models in this study.

The calculation results in this study suggest that in 2011  ${}^{134}$ Cs and  ${}^{137}$ Cs in the canopy, organic layer and top 5 cm of soil were the main contributors to the elevated ambient dose equivalent rates at 1 m above the ground in these forests in eastern Japan. By 2014 the organic layer and top 5 cm of soil were the main contributors to  $\dot{H}^*(10)$  at 1 m above the ground (excluding natural background) as by then most  ${}^{134}$ Cs and  ${}^{137}$ Cs fallout intercepted by the canopy had migrated to the forest floor. By 2016 and 2017  ${}^{134}$ Cs and  ${}^{137}$ Cs in the top 5 cm of forest soil had become the dominant contributor to  $\dot{H}^*(10)$  at 1 m.

Uncertainties in the simulation results were derived assuming the measurement 467 uncertainties in the model inputs had Gaussian distributions. The measurement uncertainties in 468 the <sup>134</sup>Cs and <sup>137</sup>Cs radioactivity of the top 5 cm of soil, the organic layer and canopy, the 469 distribution of <sup>134</sup>Cs and <sup>137</sup>Cs within 5 cm soil layers and the natural background component 470 were the main contributors to the total uncertainties in simulated  $\dot{H}^*(10)$ . Better 471 characterization of the distribution of <sup>134</sup>Cs and <sup>137</sup>Cs with mass depth in soil, and site-specific 472 evaluation of natural background contributions to  $\dot{H}^*(10)$  using in-situ gamma spectroscopy, 473 would help lower the modelling uncertainties. The former is especially important post-2017 474 due to <sup>134</sup>Cs and <sup>137</sup>Cs in the top 5 cm of soil being the main contributor to  $\dot{H}^*(10)$  at 1 m above 475 the ground. 476

# 478 Acknowledgements

The authors thank researchers in the Department of Plant Ecology, the Department of Forest Vegetation, the Department of Forest Soils and the Department of Wood Properties and Processing of FFPRI for allowing their datasets to be used for this research. We thank Dr. Kimiaki Saito for advice and comments on the manuscript. We also thank colleagues in the Center for Computational Science & e-Systems and the Environmental Research Group of CLADS within JAEA for their support. The simulations were performed on JAEA's SGI ICEX supercomputer.

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# 748 **Figure Captions**

Figure 1. Relaxation mass depths obtained from fitting the profiles for the vertical distribution
 of <sup>137</sup>Cs within soil at forests surveyed by the University of Tsukuba.

Figure 2. Schematic of the simulation model used for forests. Heights  $h_{\rm T}$  and  $h_{\rm LLB}$  represent stand density weighted mean tree heights and lowest live branch heights, respectively. Details of the organic layer and soil layers are shown in the enlarged region. Colors of layers match those used in Figs. 5–7.

Figure 3. Convergence of estimates for uncertainties in simulated  $\dot{H}^*(10)$  at 1 m height when varying different input variables in their uncertainties distribution. Data is for site OT-S in 2016. Black line showing total simulation uncertainty was calculated by varying all input variables randomly in their uncertainty distributions.

Figure 4. Scatter plots showing correlation between measured and simulated  $\dot{H}^*(10)$  in the forests at (a) 0.1 m and (b) 1 m heights above the ground surface. As guides solid lines indicate equality and dashed lines a factor of two deviation from equality are shown. Horizontal errors bars show  $\pm$  one standard error of  $\dot{H}^*(10)$  measurements. Vertical errors show  $\pm$  one standard deviation of the calculated  $\dot{H}^*(10)$  from simulations varying all input variables in the Gaussian measurement uncertainty distributions.

Figure 5. Break down of contributions to  $\dot{H}^*(10)$  at 1 m height from the simulation analysis.

Figure 6. Changes in distribution of  $^{134}$ Cs and  $^{137}$ Cs between the forest layers over time.

Figure 7. As per Figure 4, but results for  $\dot{H}^*(10)$  at 20 m above the ground.

Figure 8. Graphs showing relative measurement uncertainties in model inputs (solid dots, left

vertical axis) and the consequent relative uncertainties in simulation outputs ( $\dot{H}^*(10)$  at 1 m

- height bar plots, right vertical axis) for sites OT-Q, OT-S and TB-H, for 2011 and 2016.
- 771 Different classes of input variables are grouped per the labels on the horizontal axis. For classes
- associated with more than one input variable (e.g. <sup>134</sup>Cs and <sup>137</sup>Cs radioactivities of the three

- soil layers between 5 and 20 cm depth), the vertical line shows the range of uncertainties in the
- inputs and the solid dot indicates the mean value.





















785 Figure 5





Figure 6 









793 Figure 8

Site	Distance	Tree species groups	Mean	Survey
code	from		stand	period
	FDNPP		density	
	(km)		(ha <sup>-1</sup> )	
KU1-H	27	Japanese cypress	1330	2012-17
KU1-Q	27	Konara oak and other broadleaf	1733	2012-17
KU1-S	26	Japanese cedar / mixed broadleaf	954 / 506	2011-17
KU2-S	28	Japanese cedar / Japanese larch / mixed	725 / 199 /	2011-17
		broadleaf	33	
OT-P	66	Japanese red pine / mixed broadleaf	853 / 370	2011-16
OT-Q	66	Konara oak and other broadleaf / Japanese	543 / 563	2011-17
		red pine		
OT-S	66	Japanese cedar	1117	2011-16
TB-H	160	Japanese cypress	2063	2011-16
TD-S	134	Japanese cedar / mixed broadleaf	1090 / 132	2011-16

Table 1. Details of the nine FFPRI forest sites that were modelled. Mean stand densities are

averages over full survey periods.

Table 2. Elemental compositions of component materials used to make up simulation layers.

	Elemental composition (wt%)	Notes
Air	H, 0.064; C, 0.014; N, 75.1;	Bellamy et al. (2018), volumetric
	O, 23.6; Ar, 1.28	density 1.2 kg/m <sup>3</sup>
Dry organic matter	H, 6.04; C, 50.4; N, 0.302;	Haraguchi et al. (1985) and
(trunk and canopy	O, 43.3	Fukushima (2010)
layers)		
Dry organic matter	H, 6.05; C, 43.2; N, 2.72;	Mix cellulose and nitrogen, C/N mass
(organic & soil	O, 48.0	ratio 15.9 (Kawada, 1982; 1989)
layers)		
Dry black forest	O, 46.3; Mg, 1.69; Al, 12.6;	AIST (2018)
soil (andosol)	Si, 23.7; Ca, 2.43; Ti, 0.979;	
	Fe, 10.7	
Water (moisture)	H, 11.2; O, 88.8	Volumetric density 1000 kg/m <sup>3</sup>

Table 3. Mean physical data over all survey years for each site. Note masses per unit areainclude the moisture contents of the layer. For the canopy and trunk layers, they also include

Site	Canopy	Tree	Mass per	Mass per unit area (kg/m <sup>2</sup> )						
	base height (m)	height (m)	Canopy layer	Trunk layer	Organic layer	Soil 0– 5 cm	Soil 5– 10 cm	Soil 10– 15 cm	Soil 15– 20 cm	
KU1-S	6.7	14.2	15.4	36.6	2.9	29.1	38.4	42.0	44.8	
KU1-H	4.0	16.9	21.0	32.7	2.8	31.1	39.4	41.6	40.7	
KU1-Q	5.6	10.2	8.3	19.1	2.4	31.9	37.3	41.1	45.8	
KU2-S	11.0	17.9	21.4	65.7	5.0	21.3	29.3	36.4	37.1	
OT-S	11.9	18.1	16.8	54.7	3.6	25.4	31.5	33.3	34.3	
OT-Q	7.4	12.1	10.8	29.4	2.8	29.0	37.0	40.3	42.6	
OT-P	7.8	12.2	9.7	31.1	3.2	30.7	37.8	39.6	41.0	
TD-S	7.7	14.3	14.6	36.8	2.8	32.7	41.7	52.2	55.0	
TB-H	10.8	17.7	20.8	67.6	2.4	33.7	43.6	42.4	40.1	
Max.	11.9	18.1	21.4	67.6	5.0	33.7	43.6	52.2	55.0	
Min.	4.0	10.2	8.3	19.1	2.4	21.3	29.3	33.3	34.3	

800 the air mass (volumetric density  $1.2 \text{ kg/m}^3$ ).

801 Table 4. Correspondence between simulated and measured  $\dot{H}^*(10)$ .

Metric	Height above ground			
	0.1 m	1 m		
Root mean square deviation (µSv/h)	0.323	0.303		
Mean percentage deviation (%)	18.9	15.7		
Mean absolute percentage deviation (%)	23.6	21.6		
Person correlation coefficient	0.963	0.952		

# 803 Supplementary Information to Calculations for ambient dose equivalent rates in nine forests in eastern Japan from <sup>134</sup>Cs and <sup>137</sup>Cs

# 804 radioactivity measurements

Site	Tree species groups			Data for lower preval	Survey	
	Main group	Secondary group	Tertiary group	Secondary group	Tertiary group	period
KU1-S	Japanese cedar	Mixed broadleaf	-	KU1-Q	-	2011-17
KU2-S	Japanese cedar	Japanese larch	Mixed broadleaf	OT-P	OT-Q	2011-17
OT-P	Japanese red pine	Mixed broadleaf	-	OT-Q	-	2011-16
OT-Q	Konara oak and other broadleaf	Japanese red pine	-	OT-P	-	2011-17
TD-S	Japanese cedar	Mixed broadleaf	-	OT-Q	-	2011-16

Table S1. Data sources for lower prevalence stand groups at sites with multiple stand groups.

Survey	Mixed Forest (MF)				Young Cedar (MC)			Mature Cedar (MC)				
	Date	β	$I_{ m inv}$	Fit $R^2$	Date	β	$I_{ m inv}$	Fit $R^2$	Date	β	$I_{ m inv}$	Fit $R^2$
_		$(g/cm^2)$	$(Bq/cm^2)$			$(g/cm^2)$	$(Bq/cm^2)$			$(g/cm^2)$	$(Bq/cm^2)$	
1	1-Jul-11	1.56	3.7	0.983	2-Jul-11	1.29	2.7	0.992	30-Jun-11	2.54	26.0	0.975
2	10-Jan-12	1.81	15.9	0.997	17-Jan-12	0.32	4.3	0.962	10-Jan-12	0.73	15.7	0.997
3	27-Aug-12	0.63	12.4	0.951	28-Aug- 12	0.65	12.0	0.964	28-Aug-12	0.99	45.0	0.923
4	11-Dec-12	0.78	23.4	0.970	13-Dec-12	0.83	21.6	0.994	13-Dec-12	1.68	21.2	0.934
5	13-Jul-13	1.07	10.0	0.993	18-Jul-13	0.87	10.3	0.957	14-Jul-13	1.67	26.3	0.999
6	22-Jul-14	1.32	20.0	0.996	21-Jul-14	0.77	30.3	0.974	21-Jul-14	0.97	39.5	0.995
7	25-Jul-15	0.49	28.1	0.996	26-Jul-15	0.67	23.9	0.999	25-Jul-15	1.30	45.4	0.996
8	25-Jun-16	0.70	22.5	0.996	7-Aug-16	0.89	31.9	0.999	7-Aug-16	1.54	49.6	0.999
9	2-Jul-17	0.86	25.3	0.992	2-Jul-17	1.54	34.1	0.995	2-Jul-17	3.10	55.0	0.993

806 Table S2. Data from fits to  $^{137}$ Cs vertical distributions within soil.

Note: Moisture contents were available from surveys 2, 3, 7 and 9 for the Mixed Forest, and surveys 2, 3, 6, 7 and 9 for the Young Cedar and

808 Mature Cedar forests, to convert masses of oven-dried soil layer samples to wet masses. This data was used directly for converting the masses for

809 these surveys. For other surveys, mean moisture contents applicable for each forest were used instead.



Figure S1. Fits of cumulative exponential distributions to cumulative distributions of <sup>137</sup>Cs with
mass depth vertically in wet soil at the sites surveyed by the University of Tsukuba. Panels (a)
to (i) show results over time from surveys 1 (in June 2011) to 9 (in June 2017)